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## Early spring mesopelagic carbon remineralization and transfer efficiency in the naturally iron-fertilized Kerguelen area

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**Abstract.** We report on the zonal variability of mesopelagic particulate organic carbon remineralization and deep carbon transfer potential during the Kerguelen Ocean and Plateau compared Study 2 expedition (KEOPS 2; October–November 2011) in an area of the polar front supporting recurrent massive blooms from natural Fe fertilization. Mesopelagic carbon remineralization (MR) was assessed using the excess, non-lithogenic particulate barium ( $\text{Ba}_{\text{xs}}$ ) inventories in mesopelagic waters and compared with bacterial production (BP), surface primary production (PP) and export production (EP). Results for this early season study are compared with the results obtained during a previous study (2005; KEOPS 1) for the same area at a later stage of the phytoplankton bloom. Our results reveal the patchiness of the seasonal advancement and of the establishment of remineralization processes between the plateau (A3) and polar front sites during KEOPS 2. For the Kerguelen plateau (A3 site) we observe a similar functioning of the mesopelagic ecosystem during both seasons (spring and summer), with low and rather stable remineralization fluxes in the mesopelagic column (150–400 m). The shallow water column ( $\sim 500$  m), the lateral advection, the zooplankton grazing pressure and the pulsed nature of the particulate organic carbon (POC) trans-

fer at A3 seem to drive the extent of MR processes on the plateau. For deeper stations ( $> 2000$  m) located on the margin, inside a polar front meander, as well as in the vicinity of the polar front, east of Kerguelen, remineralization in the upper 400 m in general represents a larger part of surface carbon export. However, when considering the upper 800 m, in some cases, the entire flux of exported carbon is remineralized. In the polar front meander, where successive stations form a time series, two successive events of particle transfer were evidenced by remineralization rates: a first mesopelagic and deep transfer from a past bloom before the cruise, and a second transfer expanding at mesopelagic layers during the cruise. Regarding the deep carbon transfer efficiency, it appeared that above the plateau (A3 site) the mesopelagic remineralization was not a major barrier to the transfer of organic matter to the seafloor (close to 500 m). There, the efficiency of carbon transfer to the bottom waters ( $> 400$  m) as assessed by PP, EP and MR fluxes comparisons reached up to 87 % of the carbon exported from the upper 150 m. In contrast, at the deeper locations, mesopelagic remineralization clearly limited the transfer of carbon to depths of  $> 400$  m. For sites at the margin of the plateau (station E-4W) and the polar front (station F-L), mesopelagic remineralization even exceeded

upper 150 m export, resulting in a zero transfer efficiency to depths > 800 m. In the polar front meander (time series), the capacity of the meander to transfer carbon to depth > 800 m was highly variable (0 to 73 %). The highest carbon transfer efficiencies in the meander are furthermore coupled to intense and complete deep (> 800 m) remineralization, resulting again in a near-zero, deep (> 2000 m) carbon sequestration efficiency there.

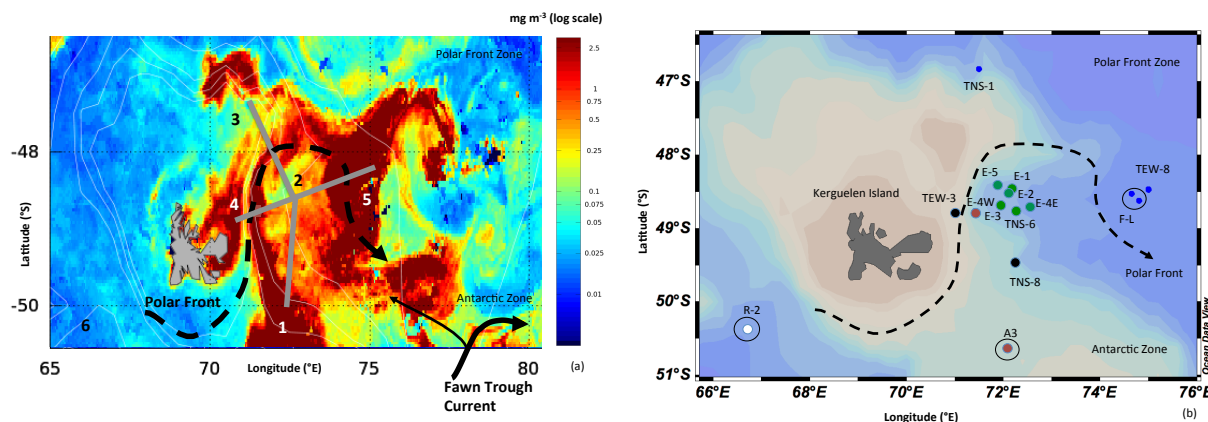
## 1 Introduction

While numerous artificial (Boyd et al., 2000, 2004; Gervais et al., 2002; Buesseler et al., 2004, 2005; de Baar et al., 2005; Hoffmann et al., 2006; Boyd et al., 2012; Smetacek et al., 2012) and natural (Blain et al., 2007; Pollard et al., 2009; Zhou et al., 2010, 2013) ocean iron-fertilization experiments in the Southern Ocean have demonstrated the role of iron in enhancing the phytoplankton biomass and production in high-nutrient low-chlorophyll (HNLC) regions, determining to what extent fertilization could modify the transfer of particulate organic carbon (POC) to the deep ocean is far from being comprehensively achieved (Lampitt et al., 2008; Morris and Charette, 2013; Le Moigne et al., 2014; Robinson et al., 2014). This is partly due to the short term over which the observations were made, precluding extrapolation to longer timescales. Moreover, when assessing whether Fe-supply could induce vertical POC transfer, the magnitude of the export from the surface is not the only important parameter to take into account. Indeed, POC fate in the mesopelagic zone (defined as the 100–1000 m depth layer) is often largely overlooked although these depth layers are responsible for the remineralization of most of the POC exported from the surface layer (Martin et al., 1987; Longhurst, 1990; Lampitt and Antia, 1997; François et al., 2002; Buesseler et al., 2007b; Buesseler and Boyd, 2009). Only few studies have considered mesopelagic carbon (C) remineralization rates (Buesseler et al., 2007a; Jacquet et al., 2008a, b, 2011a, b; Salter et al., 2007) to estimate the response of deep POC export to fertilization. Assessing mesopelagic C remineralization is pivotal for evaluating remineralization length scale as well as the timescale of the C storage in the deep ocean. Indeed the typical depth of the main thermocline, 1000 m (IPCC, WG1, 2007, chapter 5) is often referred to as the horizon, clearly removed from the surface ocean and atmosphere (Passow and Carlson, 2012). Overall, assessing mesopelagic C remineralization will allow one to better quantify the ocean's biological carbon pump and its efficiency in the global C cycle which holds large uncertainty and is currently under debate (e.g. from 5 Gt yr<sup>-1</sup> in Henson et al., 2011 to 21 Gt C yr<sup>-1</sup> in Laws et al., 2000 and 13 Gt yr<sup>-1</sup> in the IPCC WG1 report (2013, chapter 6)).

The present work aims at understanding the impact of a natural iron-induced bloom on the mesopelagic POC rem-

ineralization and zonal variability in the Kerguelen area (Southern Ocean). Here, C remineralization was assessed from particulate biogenic Ba (hereafter called excess Ba or Ba<sub>xs</sub>; mainly forms as barite BaSO<sub>4</sub> crystals) contents in the mesopelagic water column. The link between barite and C remineralization depends on the fact that this mineral precipitates in oversaturated micro-environments (biogenic aggregates) during the process of prokaryotic degradation of sinking POC (Dehairs et al., 1980, 1992, 1997, 2008; Stroobants et al., 1991; Cardinal et al., 2001, 2005; Jacquet et al., 2007, 2008b, 2011a; Planchon et al., 2013; Sternberg et al., 2007, 2008a, 2008b). Once the aggregates have been remineralized, barites are released and spread through the mesopelagic layer. Overall, earlier work highlights the fact that suspended barite in mesopelagic waters builds up over the growing season and reflects past remineralization activity integrated over several days to weeks (Dehairs et al., 1997; Cardinal et al., 2005; Jacquet et al., 2007, 2008b). An algorithm linking mesopelagic Ba<sub>xs</sub> contents to oxygen consumption (Shopova et al., 1995; Dehairs et al., 1997) allowed remineralization of POC fluxes to be estimated for the mesopelagic layer. Combined with surface C production and export estimates, mesopelagic Ba<sub>xs</sub> also highlights the deep carbon transfer efficiency of the system. From earlier studies, the efficiency of C transfer through the mesopelagic layer was reported to increase under artificially induced (EIFEX; Strass et al., 2005; Smetacek et al., 2012) and natural (KEOPS; Blain et al., 2007) Fe-replete conditions (Jacquet et al., 2008a, b; Savoye et al., 2008) compared to Fe-limited, non-bloom HNLC reference stations in the Southern Ocean. In contrast, C transfer efficiency through the mesopelagic layer was reported to be lower in natural Fe-replete locations during the SAZ-Sense (Sub-Antarctic Zone Sensitivity to Environmental Change) cruise off Tasmania (Jacquet et al., 2011a, b). Differences in plankton community structure and composition (e.g. diatoms vs. flagellates, type of diatoms) were suggested as possible causes of such discrepancies in C transfer efficiency through the mesopelagic layer (Jacquet et al., 2008a, 2011a, b). Also, differences in integration timescales for the processes that control the carbon fluxes in artificially vs. naturally Fe fertilized systems may yield an incomplete picture of the C transfer potential and lead to misleading conclusions.

Here, we examine changes in mesopelagic POC remineralization during the early spring (October–November 2011) KEOPS 2 expedition to the naturally iron-fertilized area east of Kerguelen Islands. The hydrographic structure of the Kerguelen area generates contrasted environments that are differently impacted by iron availability and mesoscale activity. The specific objectives of the present work are to assess the zonal variability of mesopelagic C remineralization and deep C transfer potential, and to identify possible causes of this variability. The same area was visited earlier in 2005 during summer at a late stage of the bloom (KEOPS 1; January–February 2005), offering a unique opportunity to estimate the main carbon fluxes over most of the growth sea-



**Figure 1.** (a) Kerguelen Island area in the Southern Ocean with KEOPS 2 sampling zones and MODIS chlorophyll concentrations ( $\text{mg m}^{-3}$ ) (Chl *a* map from 11 November 2011, courtesy of F. d'Ovidio) superimposed. 1 refers to station A3; 2 to stations E; 3 to the south–north transect; 4 to the west–east transect; 5 to station F-L and 6 to reference station R-2; (b) Corresponding stations location. Colours indicate stations with similar  $\theta - S$  and Chl *a* characteristics.

son. Mesopelagic C remineralization estimates are compared to particle and biological parameters as reported in other papers included in this issue (Cavagna et al., 2014; Christaki et al., 2014; Dehairs et al., 2014; Lasbleiz et al., 2014; Laurenceau-Cornec et al., 2015; Planchon et al., 2014; van der Merve et al., 2015) and in Blain et al. (2007), Christaki et al. (2008), Jacquet et al. (2008a), Park et al. (2008) and Savoye et al. (2008).

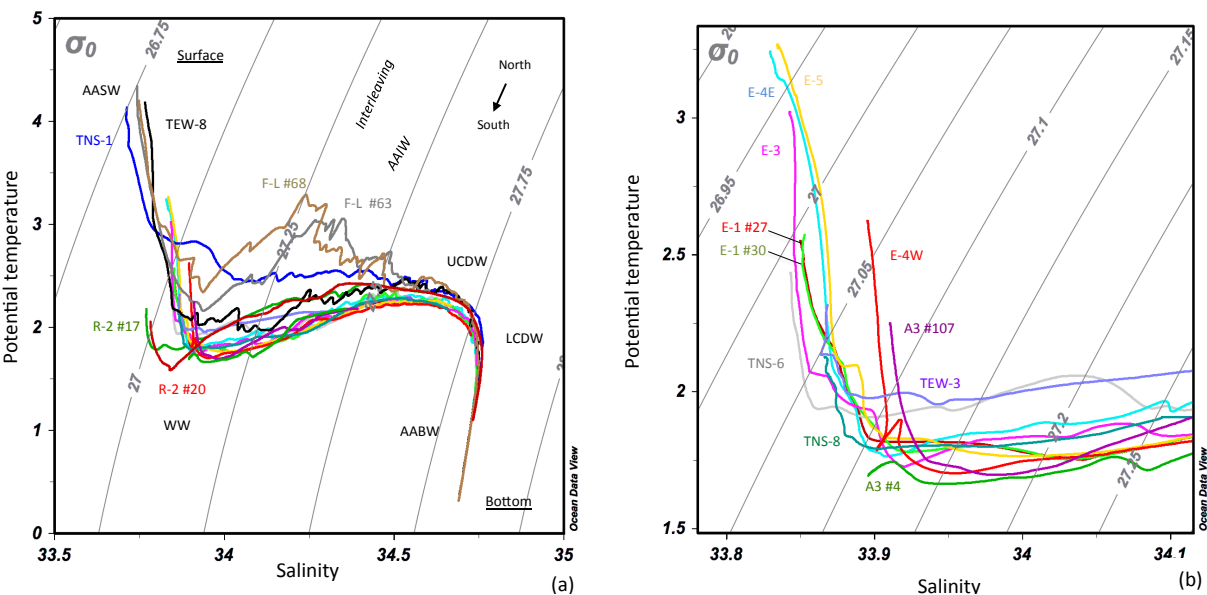
## 2 Experiment and methods

### 2.1 Study area

The KEOPS 2 (Kerguelen Ocean and Plateau compared Study) cruise was conducted in austral spring at the onset of the bloom from 10 October to 20 November aboard the R/V *Marion Dufresne* (TAAF/IPEV). The KEOPS 2 expedition studied the Kerguelen Plateau area (Indian sector of the Southern Ocean) which is characterized by the passage of the polar front (PF), as illustrated in Fig. 1a. The Kerguelen Plateau is surrounded by the Antarctic Circumpolar Current (ACC) whose main branch circulates to the north of the plateau (Park et al., 2008). A second branch of the ACC circulates to the south of Kerguelen Islands to later join a branch of the Fawn Trough Current (FTC). The FTC has a main northeast direction, but a minor branch splits away northwestward to join the eastern side of the Kerguelen Plateau (Park et al., 2008; Fig. 1a). These particular hydrographic features generate a mosaic of recurrent massive bloom patterns in the northeastern part of the plateau, and the possible sources and mechanisms for fertilization were investigated during ANTARES 3 (1995; Blain et al., 2001) and the KEOPS 1 cruise (January–February 2005, late summer conditions; Blain et al., 2007, 2008). During KEOPS 2 the evolution of Chl *a* data based on multi-satellite imagery

of the study area revealed the presence of different Chl *a* rich plumes (D'Ovidio et al., 2015) (Fig. 1a; e.g. Chl *a* map from 11 November 2011). Stations were sampled in distinct zones covering these different bloom patterns (Fig. 1a) (corresponding stations are reported in Fig. 1b): (a) on the shallow plateau (station A3; see 1 in Fig. 1a). Note that station A3 coincides with a site studied during the KEOPS 1 cruise, and that it was sampled twice over a 27-day period; (b) in a meander formed by a quasi-permanent retroflexion of the polar front (PF) and topographically steered by the eastern escarpment (Gallieni Spur) of the Kerguelen Plateau (mainly stations E, sampled as a quasi-Lagrangian temporal series) (see 2 in Fig. 1a); (c) along a north–south transect (referred to as TNS stations; see 3, grey line in Fig. 1a) and a east–west transect (referred to as TEW stations; see 4, grey line in Fig. 1a), both crossing the PF; and (d) in the polar front zone (PFZ) in the vicinity (east) of the PF (station F-L; see 5 in Fig. 1a). Furthermore we also sampled a reference HNLC, non-bloom, non-Fe-fertilized station southwest of the plateau (station R-2; see 6 in Fig. 1a). Station locations are given in Table 1.

Detailed descriptions of the complex physical structure of the area, circulation, water masses and fronts are given in Park et al. (2014). Briefly, the main hydrodynamic features observed during the cruise are the following (see  $\theta - S$  diagram, Fig. 2a): (1) north of the PF, stations in the PFZ (TNS-1, TEW-8 and F-L) present Antarctic Surface Waters (AASW;  $\theta \simeq 4^\circ\text{C}$  and density  $< 27$ );  $\theta - S$  characteristics between 150 to 400 m at station F-L (and to a lesser extent at station TNS-1) reveal the presence of interleaving with waters from northern (subantarctic) origin, centred between the 27.2 and 27.5  $\text{kg m}^{-3}$  density curves, where Antarctic Intermediate Waters (AAIW) are usually found. This contrasts with the conditions at station TEW-8, where there is no evidence of interleaving; (2) stations south of



**Figure 2.** (a) Potential temperature  $\theta$ –salinity  $S$  plots and isopycnals for KEOPS 2 profiles, (b) Focus on the upper 200 m water column. AASW: Antarctic Surface Waters, AAIW: Antarctic Intermediate Waters, WW: winter waters, UCDW and LCDW: Upper and Lower Circumpolar Deep Water, AABW: Antarctic Bottom Water. Graph constructed using Ocean Data View (Schlitzer, 2002; Ocean Data View; <http://www.awi-bremerhaven.de/GEO/ODV>).

**Table 1.** Station locations, cast numbers and bottom depths during KEOPS 2. Depth-weighted average (DWA) values of mesopelagic  $Ba_{XSS}$  (pM) and  $Ba_{XSS}$ -based mesopelagic POC remineralization (MR;  $\text{mg C m}^{-2} \text{ day}^{-1}$ ) integrated between 150–400 and 150–800 m depths. See text for further information on calculation.

Station	CTD cast no.	Long (° E)	Lat (° S)	Date of sampling	Seafloor [m]	DWA $Ba_{XSS}$ [pM] 150–400 m	DWA $Ba_{XSS}$ [pM] 150–800 m	MR 150–400 m [ $\text{mg C m}^{-2} \text{ day}^{-1}$ ]	MR Std. Uncertainty %	MR 150–800 m [ $\text{mg C m}^{-2} \text{ day}^{-1}$ ]	MR Std. Uncertainty %
Plateau											
A3-1	4 <sup>a</sup>	72.080	50.629	20 Oct 2011	530	316	–	14	4	–	–
A3-2	107 <sup>a</sup>	72.056	50.624	16 Nov 2011	527	267	–	11	5	–	–
TEW-3	38	71.018	48.799	31 Oct 2011	560	324	–	28	8	–	–
Meander time series											
TNS-6	10	72.277	48.779	22 Oct 2011	1885	427	389	31	7	69	17
E-1	27/30	72.187	48.458	29, 30 Oct 2011	2056	387	325	26	6	48	14
E-2	43	72.077	48.523	1 Nov 2011	2003	301	309	15	5	42	13
E-3	50/55	71.967	48.702	3, 4 Nov 2011	1915	258	286	10	4	35	12
E-4E	94/97	72.563	48.715	13, 14 Nov 2011	2210	395	357	27	7	58	15
E-5	113/114	71.900	48.412	18 Nov 2011	1920	402	380	28	7	66	17
Polar front zone											
TNS-1	15	71.501	46.833	23 Oct 2011	2280	350	315	22	6	45	14
TEW-8	47	74.999	48.471	2 Nov 2011	2786	199	240	2	4	20	11
F-L	63/68	74.659	48.532	6, 7 Nov 2011	2695	345	328	21	6	49	14
Polar front											
E-4W	81/87	71.425	48.765	11, 12 Nov 2011	1384	468	411	36	8	76	18
Antarctic zone											
R-2 (reference site)	17/20	66.717	50.359	25, 26 Oct 2011	2300	572	456	50	10	91	20
TNS-8	8	72.240	49.463	21 Oct 2011	1030	473	358	37	8	59	15

<sup>a</sup> Station A3 (CTD no. 4 and 107); integration up to 354 and 405 m.

the PF exhibit subsurface temperature minima characteristic of winter waters (WW); below the WW, three water masses can be identified, namely: the Upper (temperature maximum) and Lower (salinity maximum) Circumpolar Deep Water (UCDW and LCDW), and the Antarctic Bottom Water (AABW). These water masses are present roughly in the following depth intervals: 700 m < UCDW < 1500 m; 1500 m < LCDW < 2500 m; AABW > 2500 m.

Based on the  $\theta - S$  characteristics (Fig. 2a, b) and surface phytoplankton biomasses, we can schematically group the stations as follows. The R-2 HNLC reference station (white dot in Fig. 1b) is characterized by a very low biomass (with low iron contents; Qu  rou   et al., 2015). Stations TEW-3 and TNS-8 (black dots) are characterized by low to moderate biomass and Fe contents. Stations A3 and E-4W (red dots; south of the PF) as well as stations TNS-1, F-L and TEW-8 (blue dots; north of the PF) are characterized by high biomass and iron contents. Stations in the core of the PF meander (green dots; stations TNS-6, E-1, E-2, E-3, E-4E and E-5 considered as a temporal series) are characterized by moderate biomass and iron contents.

## 2.2 Sampling and analyses

Twenty-two CTD (conductivity, temperature, and depth) casts (surface to 500–2000 m) were sampled for particulate barium (Table 1) using a CTD rosette equipped with twenty-two 12 L Niskin bottles. Deep particulate Ba profiles ( $> 1000$  m) were not systematically obtained from the same CTD cast, but from successive casts sampled closely in time and space and having similar  $\theta - S$  data profiles. In the following, we use both the station and CTD numbers to refer to stations.

Four to 7 L of seawater were filtered onto 47 mm polycarbonate membranes (0.4  $\mu$ m porosity) under slight overpressure supplied by filtered air (0.4  $\mu$ m). The filters were rinsed with Milli-Q grade water ( $< 5$  mL) to remove sea salt, dried (at 50  $^{\circ}$ C) and stored in Petri dishes for later analysis. In the laboratory, we performed a total digestion of samples using a tri-acid (0.5 mL HF / 1.5 mL HCl / 1 mL HNO<sub>3</sub>; all Suprapur grade) mixture in closed Teflon beakers overnight at 90  $^{\circ}$ C in a clean pressurized room. After evaporating until nearly dry, samples were re-dissolved into around 13 mL of 2 % HNO<sub>3</sub>. The solutions were analysed for Ba and other major and minor elements using a ICP-QMS (inductively coupled plasma quadrupole mass spectrometer; X Series 2 Thermo Fisher) equipped with collision cell technology (CCT). To correct instrumental drift and matrix effects, internal standards and matrix-matched calibrations were used. We analysed several certified reference materials which consisted of dilute acid-digested rocks (BHVO-1, JB-3 and JGb-1), natural riverine water (SLRS-5) and multi-element artificial solutions for these external calibrations. Based on analyses of these external standards, accuracy and reproducibility are both better than  $\pm 5$  %. For more details on sample processing and analysis we refer to Cardinal et al. (2001). Among all elements analysed, we were particularly interested in Ba and Al. The presence of sea salt was checked by analysing Na and the sea-salt particulate Ba contribution was found to be negligible. Average detection limits equal 0.6 nM for Al and 3 pM for Ba. Detection limits were calculated as 3 times the standard deviation of the blank measured on board and then normalized to an average dilution factor of 385, i.e. particles

from around 5 L of Milli-Q water, dissolved in a final volume of 13 mL as for the samples. Biogenic barium (hereafter called excess Ba or Ba<sub>xs</sub>) was calculated as the difference between total particulate Ba and lithogenic Ba using Al as the lithogenic reference element (Dymond et al., 1992; Taylor and McLennan, 1985). At most sites and depths, the biogenic Ba<sub>xs</sub> represented  $> 95$  % of total particulate Ba. Lithogenic Ba reached up to 20 % of total particulate Ba at some depths in the upper 80–100 m mainly at station R-2 and stations north of the polar front (i.e. TEW-8, F-L and TNS-1). The standard uncertainty (Ellison et al., 2000) of Ba<sub>xs</sub> data ranges between 5 and 5.5 %. Ba<sub>xs</sub> and Al data are reported in Appendix A.

## 2.3 O<sub>2</sub> consumption and POC remineralization

The rate of oxygen consumption and particulate organic carbon remineralization rate in the mesopelagic layer (later referred to as MR) can be estimated using an algorithm relating mesopelagic Ba<sub>xs</sub> contents and oxygen consumption based on earlier observations in the Southern Ocean (Shopova et al., 1995; Dehairs et al., 1997, 2008). The detailed calculations are described in Jacquet et al. (2008a, 2011a). Briefly, we use the following equations:

$$J_{O_2} = (Ba_{xs} - Ba_{residual}) / 17\,450, \quad (1)$$

$$C_{respired} = Z \times J_{O_2} \times RR, \quad (2)$$

where  $J_{O_2}$  is the O<sub>2</sub> consumption ( $\mu$ mol L<sup>-1</sup> day<sup>-1</sup>) and  $C_{respired}$  is the mineralization rate of organic carbon (in mmol C m<sup>-2</sup> day<sup>-1</sup>; MR); Ba<sub>xs</sub> is the depth-weighted average (DWA) Ba<sub>xs</sub> value, i.e. the Ba<sub>xs</sub> inventory divided by the depth layer considered Z, Ba<sub>residual</sub> is the residual Ba<sub>xs</sub> signal (or Ba<sub>xs</sub> background) at zero oxygen consumption and RR is the Redfield C / O<sub>2</sub> molar ratio (127/175; Broecker et al., 1985). DWA Ba<sub>xs</sub> values were calculated for both layers at 150 to 400 m (plateau and deep stations) and 150 to 800 m (deep stations only; see details further below). The residual Ba<sub>xs</sub> is considered as “preformed” Ba<sub>xs</sub>, left over after partial dissolution and sedimentation of Ba<sub>xs</sub> produced during a previous phytoplankton growth event. In BaSO<sub>4</sub>, saturated waters, such as the ones filling the whole ACC water column (Monnin et al., 1999), this background Ba<sub>xs</sub> value was considered to reach 180 to 200 pM, which is rather characteristic for the deep ocean ( $> 1000$  m) (see Dehairs et al., 1997; Jacquet et al., 2008a, 2011). In the present study we used a Ba<sub>xs</sub> background of 180 pM.

We take the opportunity here to also compare O<sub>2</sub> consumption rates for the KEOPS 1 expedition (D. Lef  vre, unpublished data) with KEOPS 1 Ba<sub>xs</sub> data published earlier (Jacquet et al., 2008a). No such O<sub>2</sub> consumption data are available for KEOPS 2. During KEOPS 1, dark community respiration (DCR) was estimated from changes in the dissolved oxygen concentration over 72 h incubations. Discrete samples were collected at three depths in the mesopelagic

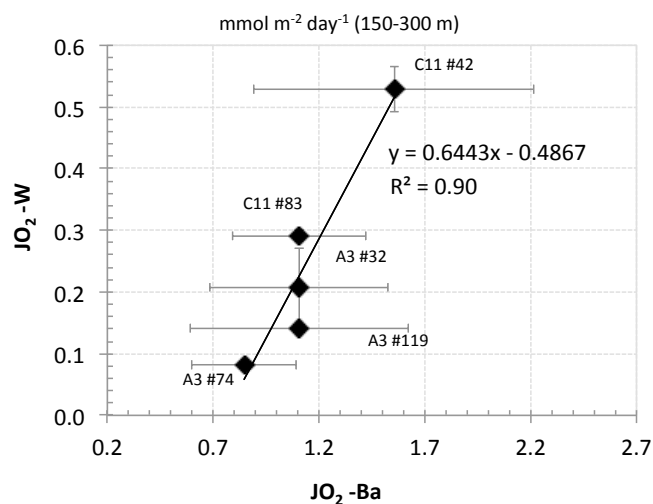
zone from Niskin bottles and placed into 125 cm<sup>3</sup> borosilicate glass bottles according to the WOCE (World Ocean Circulation Experiment) procedure, and oxygen concentration was determined by Winkler titrations using a photometric endpoint detector (Williams and Jenkinson, 1982). By integrating DCR data with the water column, we estimated the rate of oxygen consumption (referred to as JO<sub>2</sub>-W). We compared JO<sub>2</sub>-W obtained from incubated oxygen samples with the rate of oxygen consumption based on KEOPS 1 mesopelagic Ba<sub>xs</sub> contents (Eq. 1; later referred to as JO<sub>2</sub>-Ba). Dissolved oxygen was measured three times at station A3 (same location as during KEOPS2) over a 19-day period (A3 CTD no. 32, 74 and 119). Dissolved oxygen was also measured at station C11, located off-shelf in less productive HNLC waters (51.65° S, 78.00° E; not shown in Fig. 1) and was sampled two times over a 10-day period (C11 CTD no. 42 and 83). Figure 3 compares JO<sub>2</sub>-W and JO<sub>2</sub>-Ba for repeat stations A3 (no. 32, 74 and 119) and C11 (no. 42 and 83) (integration between 150–300 m). JO<sub>2</sub>-W rates range from 0.082 to 0.208 mmol m<sup>-2</sup> day<sup>-1</sup> at station A3 and from 0.292 to 0.528 mmol m<sup>-2</sup> day<sup>-1</sup> at station C11. Although JO<sub>2</sub>-Ba rates (from 0.846 to 1.555 mmol m<sup>-2</sup> day<sup>-1</sup>) are slightly higher than JO<sub>2</sub>-W, JO<sub>2</sub> rates are of the same order of magnitude and present the same trend. We observed a significant positive correlation between the JO<sub>2</sub> rates ( $R^2 = 0.90$ ;  $p < 0.01$ ), with a slope of 0.64. The difference in oxygen consumption rates can be explained by the integration time of both methods (a few hours for the incubations vs. few days to weeks for Ba<sub>xs</sub>) and by the fact that KEOPS 1 occurred at the decline of the bloom (late summer; low organic substrates), which would explain the lower JO<sub>2</sub> rates as estimated by the incubation method.

Overall, these results highlight the need for further constraining spatial and temporal variability of deep ocean oxygen utilization via a combination of direct rate measurements and the Ba<sub>xs</sub> proxy. In the present work, O<sub>2</sub> consumption and POC remineralization were assessed from Ba<sub>xs</sub> inventories and Eqs. (1) and (2). C remineralization rates are given in Table 1. Relative standard uncertainties (Ellison et al., 2000) of C remineralization ranged between 4 and 20 %.

### 3 Results

#### 3.1 Particulate biogenic Ba<sub>xs</sub> profiles

Ba<sub>xs</sub> profiles in the upper 800 m are reported in Fig. 4. The complete whole water column data set is given in Appendix A. From previous studies we know that Ba<sub>xs</sub> in surface waters is distributed over different, mainly non-barite biogenic phases (see Stroobants et al., 1991; Jacquet et al., 2007; Cardinal et al., 2005; Sternberg et al., 2005). As such, these do not reflect POC remineralization processes, in contrast to mesopelagic waters where Ba<sub>xs</sub> is mainly composed of barite (Dehairs et al., 1980) formed during prokaryotic

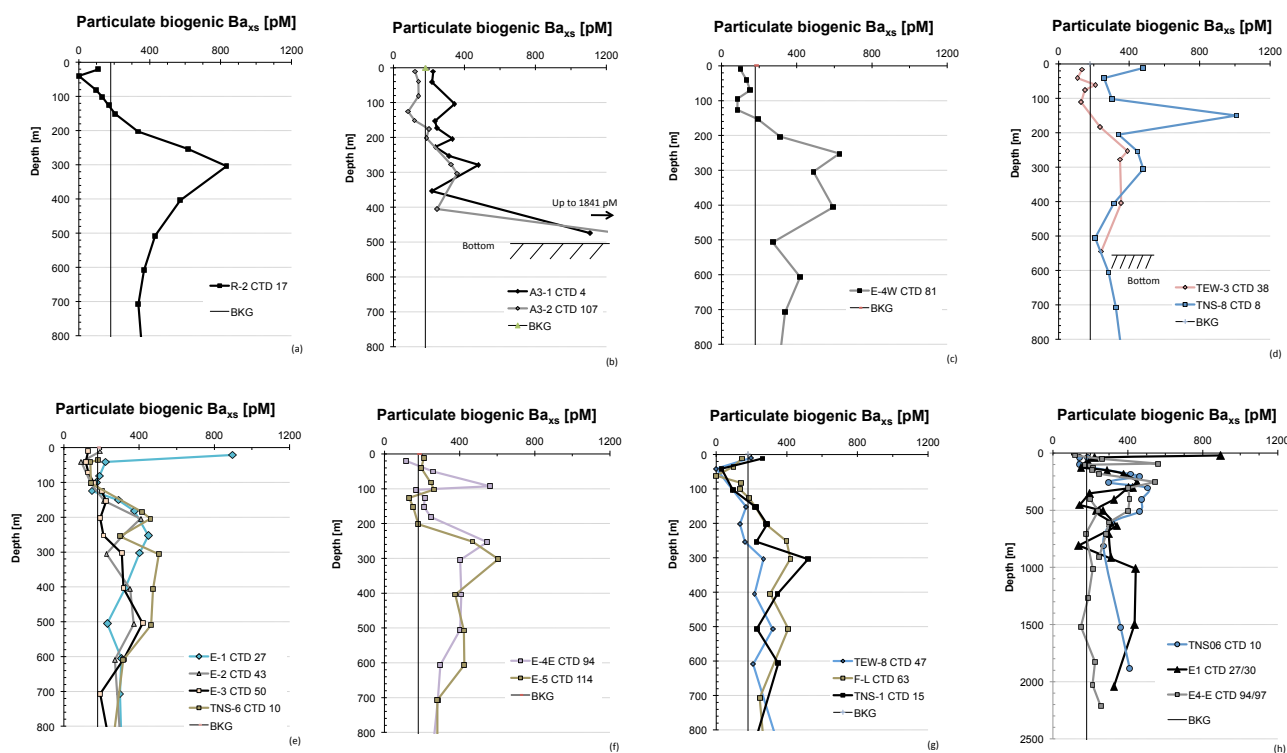


**Figure 3.** Rates of oxygen consumption (mmol m<sup>-2</sup> day<sup>-1</sup>) during KEOPS 1 as directly measured (JO<sub>2</sub>-W) and from mesopelagic Ba<sub>xs</sub> contents (JO<sub>2</sub>-Ba). Rates are integrated between 150–300 m.

degradation of sinking POC (Martin et al., 1987; Sarmiento et al., 1993; Buesseler et al., 2007b). For KEOPS 2 we observed that Ba<sub>xs</sub> concentrations generally increase below 150 m (i.e. they increase above the background level set at 180 pM), but some sites have ocean surface Ba<sub>xs</sub> contents significantly larger than background (E-1, 896 pM at 21 m; E4-E, 563 pM at 93 m). Such values are not unusual, and very high surface values have been observed occasionally in earlier Southern Ocean studies. During KEOPS 1, surface Ba<sub>xs</sub> maxima at the three A3 repeat stations ranged from 1354 to 5930 pM at 50 m, likely associated with phytoplankton-derived particles (Jacquet et al., 2008a).

The following part focuses on the mesopelagic zone where most of the remineralization of exported organic matter takes place. The Ba<sub>xs</sub> profile for station R-2 (CTD no. 17) displayed a characteristic mesopelagic Ba<sub>xs</sub> maximum, reaching up to 834 pM at 304 m, which is actually one of the highest values observed for the whole study (Fig. 4a). Ba<sub>xs</sub> profiles for stations A3 above the Kerguelen Plateau (A3-1 CTD no. 4 and A3-2 CTD no. 107; Fig. 4b) had lower mesopelagic Ba<sub>xs</sub> content, with values ranging from about 80 to 350 pM. For both A3 visits, Ba<sub>xs</sub> values increased close to the seafloor, reaching up to 1108 pM (A3-1, 474 m) and 1842 pM (A3-2, 513 m). In contrast, station E-4W (located further north along the margin in deeper waters, but with similar  $\theta - S$  and Chl *a* characteristics as station A3) displayed a large mesopelagic Ba<sub>xs</sub> maximum reaching up to 627 pM at 252 m (Fig. 4c). Station TEW-3 (located on the Kerguelen Plateau, in waters with similar  $\theta - S$  and Chl *a* characteristics as station TNS-8) had a profile similar to the one observed at station A3-2, but compared to plateau sites A3-1 and A3-2, no increased Ba<sub>xs</sub> contents were observed in bottom water (Fig. 4d). The other stations of the study area (Fig. 4d–g) have Ba<sub>xs</sub> pro-





**Figure 4.** Particulate biogenic  $Ba_{xs}$  profiles (pM) in the upper 800 m (Fig. 4a–g) and in the upper 2500 m (Fig. 4h). Stations are identified by CTD cast numbers. BKG:  $Ba_{xs}$  background (180 pM).

files similar to the one at station E-4W, showing the characteristic  $Ba_{xs}$  maximum between 200 and 500 m. Note that for most of the stations,  $Ba_{xs}$  concentrations in waters below the mesopelagic maximum did not systematically decrease to reach the  $Ba_{xs}$  background level (180 pM; see above). In some cases  $Ba_{xs}$  contents significantly higher than residual  $Ba_{xs}$  were observed until below 1000 m (see Appendix A). This is particularly salient at stations TNS-6, E-1, E-2 and F-L where  $Ba_{xs}$  values below 1000 m reach 410 pM at 1886 m (TNS-6) and 436 pM at 1498 m (E-1). These cases of high, deep  $Ba_{xs}$  contents clearly contrasted with the values observed at station E4-E (Fig. 4h).

### 3.2 Depth-weighted average $Ba_{xs}$ content of mesopelagic waters

Since the base of the mixed layer was generally shallower than 150 m, this depth is taken as the upper boundary of the mesopelagic domain. The depth-weighted average (DWA)  $Ba_{xs}$  contents, calculated for the 150–400 and 150–800 m depth intervals, are given in Table 1. For the profiles on the plateau (500 m water column), bottom waters with evidence of sediment resuspension were not taken into account when calculating DWA  $Ba_{xs}$  values ( $\geq 400$  m). Particle size spectra indicated that sediment resuspension occurred especially at stations A3 and TEW-3 (Jouandet et al., 2014; Lasbleiz et al., 2014; van der Merve et al., 2015). Thus, at site A3

(Fig. 4b), DWA  $Ba_{xs}$  was calculated for the layer between 150 and 354 m for A3-1 (CTD no. 4) and between 150 and 405 m for A3-2 (CTD no. 107). For station TEW-3 (CTD no. 38), DWA  $Ba_{xs}$  was calculated for the water layer between 150 and 400 m (Fig. 4d). For the deep sites, we considered both the 150–400 and the 150–800 m depth intervals when calculating the DWA  $Ba_{xs}$  contents. Depth-weighted average  $Ba_{xs}$  values were translated into carbon remineralization rates using Eqs. (1) and (2) given above. These rates ranged from 2 to 91 mg C m<sup>-2</sup> day<sup>-1</sup> (Table 1).

DWA  $Ba_{xs}$  values range from 199 to 572 pM (Table 1) and fit within the range reported for polar front areas during previous studies (Cardinal et al., 2001, 2005; Jacquet et al., 2005, 2008a, b, 2011; Planchon et al., 2013). For the KEOPS 2 cruise the main observed features are as follows:

- Unexpectedly, the highest DWA  $Ba_{xs}$  value of the whole study area (572 pM; 150–400 m) was observed at the reference R-2 site. For R-2, Bowie et al. (2014), Qu  rou   et al. (2015) and van der Merve et al. (2015) reported local maxima in particulate and dissolved trace metals at 500 m and deeper, reflecting lateral transport of lithogenic matter possibly originating from the Leclaire Rise (a large seamount located west of R-2). Similarly, Lasbleiz et al. (2014) observed a maximum of lithogenic silica (LSi) at 500 m, confirming lithogenic inputs there. However, we note that the



mesopelagic  $Ba_{xs}$  maximum at R-2 occurs at shallower depths, around 300 m, and that there is no evidence for elevated values at 500 m where the previous authors reported higher trace element and silica concentrations. Also, as reported above (see Sect. 2.2 and Appendix A), the higher lithogenic Ba fractions at R-2 (up to 20 % of the total Ba) occur only in the upper 80 m. Moreover, we note that surface waters at R-2 has already experienced some nitrate consumption as compared to subsurface winter waters ( $T_{min}$  waters). Indeed, surface waters had 10 % less nitrate than winter water (26  $\mu$ M at 5 m vs. 29  $\mu$ M at 200 m), and the isotopic enrichment of this surface nitrate confirmed a suggestion of uptake (see Dehairs et al., 2014). Also, Lasbleiz et al. (2014) reported relatively low Si:C and Si:N ratios for surface ocean suspended matter, pointing to the development of a diatom assemblage just prior to the sampling, consistent with the high dissolution rates of biogenic silica (BSi) that Closset et al. (2014) reported for R-2 surface waters. It is therefore likely that the mesopelagic  $Ba_{xs}$  content at R-2 indeed reflects remineralization of organic material that was fuelled by an important past early spring production and export event. Similarly, during late winter (November 1993) F. Dehairs (unpublished results) observed the presence of significant numbers of barite microcrystals in mesopelagic waters at the KERFIX time series station (50°40' S, 68°25' E) located east of R-2. Results would thus suggest the occurrence in this HNLC area of recurrent brief early spring diatom productive period pulses and subsequent export and remineralization activity in the underlying layers. Chl *a* satellite images (Giovanni – Interactive Visualization and Analysis, NASA GES DISC) corroborate that the R-2 and KERFIX area is occasionally subject to enhanced biomass during early spring;

- b. The two successive visits (27 days apart) at site A3 yielded relatively low DWA  $Ba_{xs}$  values of 267 and 316 pM, and a quite similar value was observed for the shallow station TEW-3 (324 pM), located further north on the plateau and north of the PF. Note that for comparison purposes, we recalculated the DWA  $Ba_{xs}$  and MR values of KEOPS 1 by considering upper and lower mesopelagic layer boundaries of 150 and 400 m rather than 125 and 450 m, as in Jacquet et al. (2008a). Also, in the aforementioned study the high  $Ba_{xs}$  contents observed near the seafloor were not excluded from the calculations, while they are here. These increased benthic boundary layer  $Ba_{xs}$  contents (observed also during KEOPS 2) are due to sediment resuspension which extended up to 70 m above the seafloor during KEOPS 1 (Blain et al., 2008; Venchiarutti et al., 2008; Armand et al., 2008). Because of these slightly different depth intervals over which  $Ba_{xs}$  values were integrated, the KEOPS 1 values discussed here will be slightly differ-

ent from those reported in Jacquet et al. (2008a). At the other depths, the lithogenic Ba contribution at A3 (KEOPS 2) was only minor;

- c. The time series stations in the polar front meander had DWA  $Ba_{xs}$  contents ranging from 258 to 427 pM (150–400 m), so reaching values exceeding those on the plateau. For these time series, stations' values decreased between day 0 (TNS-6) and 12 (E-3), and then increased again at days 22 (E-4E) and 27 (E-5). Stations E-4W and TNS-8, above the plateau but in deeper waters close to the Kerguelen margin, at the edge the high biomass plume (Fig. 1), had the highest DWA  $Ba_{xs}$  values (468 and 473 pM, respectively; 150–400 m), not considering the R-2 reference station. The polar front F-L site, although located within the eastern part of the high biomass plume, had a smaller DWA  $Ba_{xs}$  value of 345 pM (150–400 m) and the nearby station TEW-8 had the lowest DWA  $Ba_{xs}$  value of the whole study area (199 pM; 150–400 m).

## 4 Discussion

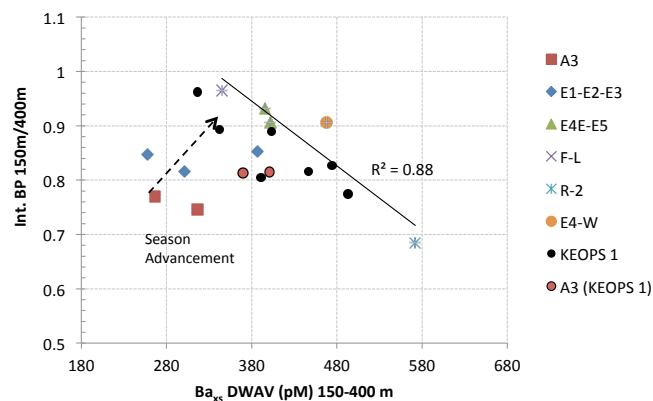
### 4.1 Mesopelagic $Ba_{xs}$ and bacterial production

Previous studies revealed that the shape of the column-integrated bacterial production (BP) profile (i.e. the attenuation length scale) was important in setting the  $Ba_{xs}$  signal in the mesopelagic zone (Dehairs et al., 2008; Jacquet et al., 2008a, 2011a). Mesopelagic  $Ba_{xs}$  content is smaller when most of the column-integrated BP is restricted to the upper mixed layer (indicating an efficient, near-complete remineralization within the surface), compared to situations where a significant part of integrated BP was located deeper in the water column (reflecting significant deep bacterial activity and POC export). During KEOPS 2 the incorporation of  $^3$ H-leucine was used to estimate bacterial production. BP data are described in Christaki et al. (2014). In Fig. 5 we compare column-integrated BP at 150 m over 400 m (BP150/400) and DWA  $Ba_{xs}$  for the 150–400 m depth interval, along with the relationship obtained during KEOPS 1 (BP200/125 and 150–450 m DWA  $Ba_{xs}$ ; Jacquet et al., 2008a; Christaki et al., 2008). Excluding stations A3, E-1, E-2 and E-3, KEOPS 2 data presented a significant correlation ( $R^2=0.88$ ;  $p<0.01$ ) and a similar trend to the one reported for KEOPS 1. A similar picture was obtained when integrating DWA  $Ba_{xs}$  and BP up to 800 m (not shown). The time series “E” stations in the meander revealed a shift from stations E-1, E-2 and E-3 to stations E-4E and E-5, i.e. towards the trend reported above. A shift was also apparent at station A3 from KEOPS 2 (early spring) to KEOPS 1 (late summer). It is thus possible that results reflect the occurrence of different stages of bloom advancement. The large variability of the  $Ba_{xs}$  and BP relationship during KEOPS 2, especially at A3 site and in the me-

**Table 2.** Comparison of mesopelagic POC remineralization (MR) with primary production (PP) and export production (EP). All fluxes in  $\text{mg C m}^{-2} \text{ day}^{-1}$ .  $r$  ratio is the ratio of Mr over EP. The C sequestration (or transfer) efficiency at 400 and 800 m (T400, T800) is the fraction of C export (EP) at 150 m exiting through the 400 m (T400) or the 800 m (T800) horizons. See text for further information on calculation.

Station	CTD	MLD [m]	Ez <sup>b</sup> [m]	PP <sup>c</sup> Ez [ $\text{mg C m}^{-2} \text{ day}^{-1}$ ]	EP <sup>d</sup> 150 m [ $\text{mg C m}^{-2} \text{ day}^{-1}$ ]	MR 150–400 m [ $\text{mg C m}^{-2} \text{ day}^{-1}$ ]	MR 150–800 m [ $\text{mg C m}^{-2} \text{ day}^{-1}$ ]	EP/PP	$r$ ratio 150–400 m	$r$ ratio 150–800 m	T400	T800
Plateau												
A3-1	4 <sup>a</sup>	161	–	–	47	14	–	–	0.29	–	0.70	–
A3-2	107 <sup>a</sup>	165	38	2172	85	11	–	0.04	0.13	–	0.87	–
Reference site												
R-2	17/20	111	92	132	30	50	91	0.23	1.65	3.02	0	0
Meander time series												
E-1	27/30	84	64	578	100	26	48	0.17	0.26	0.48	0.74	0.52
E-3	50/55	41	68	748	130	10	35	0.17	0.08	0.27	0.92	0.73
E-4E	94/97	77	34	1037	48	27	58	0.05	0.57	1.21	0.43	0
E-5	113/114	36	54	1064	84	28	66	0.08	0.33	0.78	0.67	0.22
Polar front zone												
F-L	63/68	21	29	3380	43	21	49	0.01	0.48	1.13	0.52	0
Polar front												
E-4W	81/87	67	31	3287	54	36	76	0.02	0.67	1.41	0.33	0

<sup>a</sup> Station A3 (CTD4 and 107); MR integrated up to 354 and 405 m. <sup>b</sup> Ez euphotic layer (until 1 % PAR level). <sup>c</sup> PP data from Cavagna et al. (2015). <sup>d</sup> EP data from Planchon et al. (2015).



**Figure 5.** Regression of the ratio of integrated bacterial production (BP) in the upper 150 m over integrated BP in the upper 400 m versus depth-weighted average (DWA) mesopelagic  $\text{Ba}_{\text{xs}}$  (pM; 150–400 m) during KEOPS 2. KEOPS 1 data (dots) are reported for comparison (Christaki et al., 2008; Jacquet et al., 2008a).

ander, could reflect the temporal evolution and patchiness of the establishment of mesopelagic remineralization processes in this polar front area.

#### 4.2 Fate of exported organic C in the mesopelagic zone and deep water column

An important question relates to the fate of the exported POC: how much of this POC is respired in the mesopelagic waters and how much escapes remineralization and is exported to deeper layers where longer-term sequestration is likely (see e.g. Passow and Carlson, 2012; Robinson et

al., 2014; Schneider et al., 2008). To address these questions, we defined two ratios: (1) the mesopelagic C remineralization efficiency ( $r$  ratio in Table 2), which is the ratio of mesopelagic C remineralization (MR, based on the DWA  $\text{Ba}_{\text{xs}}$  concentrations) over C export (EP) from the 150 m horizon (based on  $^{234}\text{Th}$ , see Planchon et al., 2014), and (2) the C transfer efficiency at 400 and 800 m (i.e. T400, T800 in Table 2), which is the fraction of C export (EP) at 150 m passing through the 400 m (T400) or the 800 m (T800) horizons (e.g.  $\text{T400} = \text{EP}_{400} / \text{EP}_{150} = 1 - (\text{MR} / \text{EP}_{150})$ , with  $\text{MR} / \text{EP}_{150} = r$  ratio; see above). This approach is similar to the one developed by Buesseler and Boyd (2009) stating that a conventional curve-fitting of particle flux data (i.e. power law or exponential) skews our interpretation of the mesopelagic processes. They recommended the use of combined metrics to capture and compare differences in flux attenuation. In the following, we compare MR fluxes for the different KEOPS 2 areas (reference site; plateau sites; polar front and polar front meander) and discuss remineralization and transfer efficiencies for those sites for which MR, primary production (PP) and/or EP data (Table 2) were available. PP data were estimated from uptake experiments including 24 h incubations at different PAR levels over the euphotic layer, i.e. up to the 0.01 % PAR level (Cavagna et al., 2014). EP data were estimated from  $^{234}\text{Th}$  activities and  $^{234}\text{Th} / \text{POC}$  ratios and are discussed in Planchon et al. (2014). The thorium method integrates POC export over a 1-month period ( $^{234}\text{Th}$  half live is 24.1 days). We remind the reader here that MR fluxes based on mesopelagic  $\text{Ba}_{\text{xs}}$  reflect past remineralization activity integrated over several days to a few weeks (Dehairs et al., 1997; Cardinal et al., 2005; Jacquet et al., 2007, 2008b). In order to compare

EP with MR ( $r$  ratio and transfer efficiency), we consider EP fluxes from 150 m. Results are compared with late summer KEOPS 1 results. For KEOPS 1, PP data are detailed in Lefèvre et al. (2008) and Mosseri et al. (2008), EP data are detailed in Savoye et al. (2008) and  $Ba_{xs}$  data are described in Jacquet et al. (2008a).

#### 4.2.1 Reference station R-2

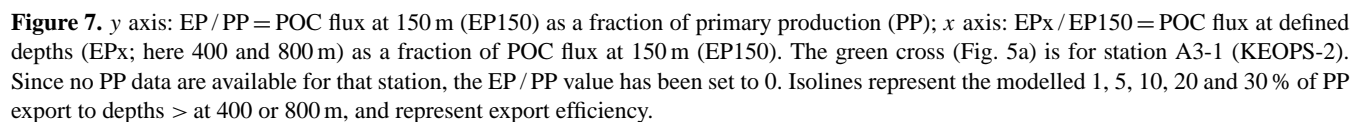
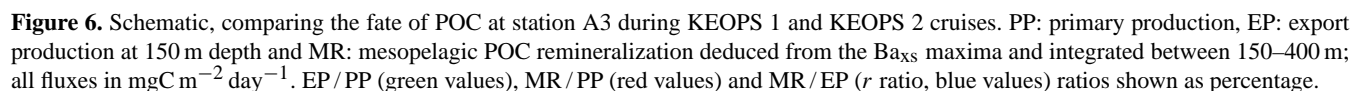
Since station R-2 had the highest DWA  $Ba_{xs}$  content, it yielded the highest MR flux of the whole study area ( $91 \text{ mg C m}^{-2} \text{ day}^{-1}$ ; 150–800 m; Table 2). In contrast, both PP and EP fluxes at R-2 were very low ( $132$  and  $10 \text{ mg C m}^{-2} \text{ day}^{-1}$ , respectively) and the calculated MR flux exceeded EP (Table 2). The resulting export efficiency (EP/PP) was high, and T400 and T800 values (the fraction of EP exported deeper than 400 and 800 m, as defined above) equal 0 (i.e. no export of POC beyond 400 and 800 m; note that  $> 100\%$  values, i.e.  $MR > EP$ , were set to zero in Fig. 7a and Table 2). The fact that MR exceeds EP therefore implies a non-steady state condition at the R-2 site. As reported above, R-2 probably experienced a brief early spring diatom production pulse days to a few weeks before the start of the KEOPS 2 cruise, followed by subsequent export and very important remineralization activity in the underlying layers as depicted by MR data.

#### 4.2.2 Station A3 on the plateau

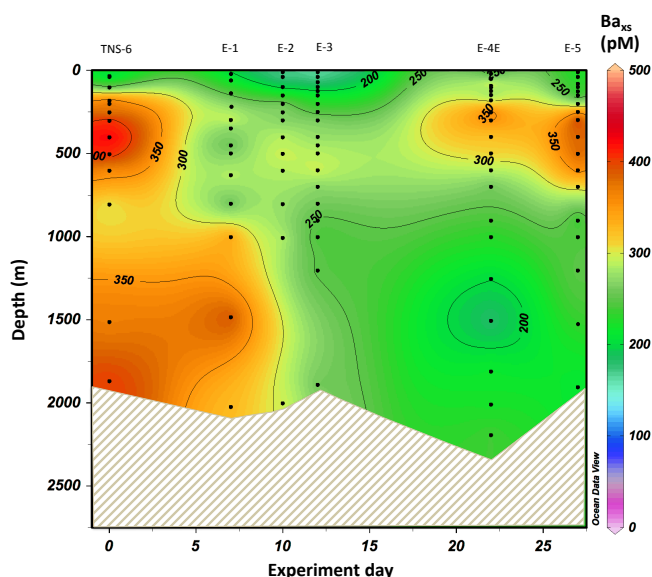
The MR fluxes on the plateau varied little between the two visits 27 days apart (Table 1) and, moreover, as discussed below they were similar to summer values obtained during KEOPS 1 (see Jacquet et al., 2008a) when the same A3 site was sampled three times over a 19-day period. While during KEOPS 2 (spring) MR fluxes at A3 ranged from  $11$  to  $14 \text{ mg C m}^{-2} \text{ day}^{-1}$  (with a standard uncertainty of around  $5\%$ ), they were slightly larger during KEOPS 1 (summer;  $17$  to  $23 \text{ mg C m}^{-2} \text{ day}^{-1}$ ) (Fig. 5). We observed differences in the mesopelagic POC remineralization efficiency between the two seasons ( $r$  ratio, blue values in Fig. 6, Table 2). During KEOPS 1,  $r$  ratios (MR/EP) remained low, ranging from  $7$  to  $9\%$  of EP at A3, while during KEOPS 2,  $r$  ratios were slightly higher but decreased from  $29\%$  (A3-1; first visit) to  $13\%$ , 27 days later (A3-2; second visit). This variation in  $r$  ratio during KEOPS 2 is mostly due to an increase of EP (from  $47$  to  $85 \text{ mg C m}^{-2} \text{ day}^{-1}$ ; Planchon et al., 2014) over the same period while MR showed little change. Although at this early stage of the season (spring) PP at A3-2 had already reached  $2172 \text{ mg C m}^{-2} \text{ day}^{-1}$  (Cavagna et al., 2014), EP remained relatively low ( $85 \text{ mg C m}^{-2} \text{ day}^{-1}$ ). Here EP accounted for only about  $4\%$  of PP (low export efficiency; see green data points in Fig. 5). These conditions suggested that phytoplankton biomass had accumulated in the surface waters without significant export at that point, or that C had been channelled to higher trophic levels as

suggested by Christaki et al. (2014). Note that a negative relationship between primary productivity and surface carbon export efficiency has already been reported from previous studies in the Southern Ocean (Lam et al., 2007; Morris et al., 2007; Savoye et al., 2008; Jacquet et al., 2011a, b). Among possible explanations for the occurrence of high-productivity low export efficiency regimes in high-latitude systems Maiti et al. (2013) mentioned differences in trophic structure, grazing intensity, recycling efficiency, high bacterial activity or increase in DOC export, but the exact reason remain unclear. In contrast, during KEOPS 1 (summer), EP fluxes reached  $250 \text{ mg C m}^{-2} \text{ day}^{-1}$  at  $125 \text{ m}$  ( $14$ – $31\%$  of PP), while PP ranged from  $865$  to  $1872 \text{ mg C m}^{-2} \text{ day}^{-1}$ , reflecting enhanced export efficiency (Jacquet et al., 2008a; Savoye et al., 2008).

It is important to underline the fact that MR at station A3 was only slightly higher in summer than in spring, especially considering the large differences in export efficiency between seasons. According to results from sediment traps deployed over 1 year at the A3 site, Rembauville et al. (2014) reported that  $60\%$  of the annual POC export at the base of the mixed layer occurred over a short periods of time representing  $< 4\%$  of the year and was composed of small highly silicified, fast-sinking, resting spores of diatoms that bypass grazing pressure. According to these authors, the pulses are linked to nutrient depletion dynamics inducing resting spore formation. During the rest of the year, the flux was composed of small diatoms (empty frustules) and small fecal pellets, with efficient C retention in the surface layer or transfer to trophic levels. If we consider that export conditions during KEOPS 2 are more similar to those prevailing most of the year, it is surprising that during KEOPS 1 (which would reflect an export event toward the end of the growth season) MR is not more important. This would indicate that fast-sinking, highly silicified and pulsed material was directly transferred to the bottom without major remineralization. Note for example that at the complex R-2 reference station, a small export event (Laurenceau-Cornec et al., 2015) held heavily silicified diatoms (Lasbleiz et al., 2014) and that the material was efficiently remineralized in the upper mesopelagic layer as witnessed by the high MR values we observed for that station. For the KEOPS 2 A3 site, Laurenceau-Cornec et al. (2015) reported that the sinking flux collected in the upper layer using gel-filled sediment traps was composed of phytodetrital aggregates that held slightly silicified diatoms (Lasbleiz et al., 2014). Even considering the shift from slightly to highly silicified material transfer between spring (KEOPS2) and summer (KEOPS 1), MR only slightly increases between both periods. Also, the mesozooplankton biomass at A3-2 was one of the highest of the KEOPS2 cruise, with a doubling from KEOPS 2 (early spring) to KEOPS 1 (late summer) (Carlotti et al., 2015). It is thus possible that at A3 the export event reported above, combined with a lasting grazing pressure, could have induced this rather low and enduring mesopelagic remineralization.



In Fig. 7a, the ratio of EP to PP (export efficiency) vs. the fraction of EP exported deeper than 400 m (i.e. T400; defined above) is shown for both KEOPS cruises. Note that for station A3-1 (KEOPS 2), there are no PP data. The A3 site shows increasing EP/PP ratios from spring (KEOPS 2) to late summer (KEOPS 1), and so do the T400 values (A3-1: 70 %; A3-2: 87 %; KEOPS 1 A3 site:  $92 \pm 1$  %). Station E-4W is located in waters with similar  $\theta - S$  and Chl  $a$  characteristics as the A3 plateau site but has a deeper water column (1384 m has PP and EP fluxes of the same order of magnitude (Table 2)). However, MR values ( $36 \text{ mg C m}^{-2} \text{ day}^{-1}$ ; 150–400 m) are larger at E-4W, resulting in a lower T400 value of around 33 %, compared to 87 % for A3-2 (Fig. 7a). When integrating between 150 and 800 m, T800 at E-4W equals 0 (i.e. no export of POC beyond 800 m; Fig. 7a and Table 2).



**Figure 8.** Temporal evolution of particulate biogenic Ba ( $Ba_{xs}$ ; pM) in the upper 2000 m water column in the polar front meander. Graph constructed using Ocean Data View (Schlitzer, 2002; Ocean Data View; <http://www.awi-bremerhaven.de/GEO/ODV>).

Station F-L (in the vicinity of the PF;  $74.7^{\circ}$  E) appears to function in a similar way as observed for E-4W ( $71.4^{\circ}$  E). PP at station F-L is relatively high ( $3380 \text{ mg C m}^{-2} \text{ day}^{-1}$ ), while EP is quite low ( $43 \text{ mg C m}^{-2} \text{ day}^{-1}$ ), reflecting the fact that the biomass was not yet exported from the surface waters or was transported to higher trophic levels. Since MR fluxes are slightly lower ( $21 \text{ mg C m}^{-2} \text{ day}^{-1}$ ; 150–400 m) at F-L than at E-4W, resulting T400 values are higher (52 %) there.

Overall, during KEOPS 2, it appears that biomass at stations A3, E-4W and F-L (sites of high productivity) had been accumulating in surface waters (e.g. transfer to higher trophic levels) and export had not yet started, considering the early stage of the season during KEOPS 2. Our observations allow us to conclude the following:

1. Both seasons (KEOPS 1 and KEOPS 2) showed a similar functioning of the mesopelagic ecosystem at A3. The rather low and enduring MR fluxes under high production and variable export regimes (high export efficiency during KEOPS 1 and low export efficiency during KEOPS 2) indicated that here mesopelagic remineralization does not represent a major resistance to organic matter transfer to the seafloor at A3. On average (considering both seasons, but excluding A3-1), the C transfer efficiency into the deep ( $> 400 \text{ m}$ ) as assessed by PP, EP and MR fluxes comparisons reached  $91 \pm 3 \%$  at A3;
2. In contrast to A3, E-4W and F-L showed important mesopelagic remineralization rates, reducing the effi-

ciency of C transfer beyond 400 m to 33 and 52 %, respectively, and to zero for both stations beyond 800 m. Bottom depth, lateral advection, zooplankton grazing pressure and the pulsed nature of the POC transfer at A3 were the particular conditions that could drive the differences in C transfer efficiency between A3 and E4-W and F-L and limit the extent of MR processes at A3.

### 4.3 Stations in the meander

Temporal short-term changes for the stations TNS-6, E-1, E-2, E-3, E-4E and E-5, located in the polar front meander, will be discussed in this section. Note that no PP or EP data exist for TNS-6. From Table 2 it appears that PP almost doubled between E-1 and E-5, but this increase was not paralleled by an increase of EP and MR, except for the 30 % EP increase from E-1 to E-3. In fact, overall EP shows a decreasing trend with time, while MR (150–400 m) stays rather constant, except for the decrease between E-1 and E-3 (Table 2). As reported above such a mismatch may result from differences in timescales characterizing the different processes that were compared. The most likely explanation is that in this early stage of the growth season, phytoplankton biomass accumulated in the surface layer and export lagged behind.

The ratio of EP to PP vs. T400 and T800 showed a large variability in transfer efficiency inside the meander (Fig. 7b). PP and EP fluxes increased by about 30 % from E-1 to E-3, but a concomitant decrease of mesopelagic MR yielded to an enhanced transfer efficiency, from 74 to 92 %, through the 400 m boundary, and from 52 to 73 % through the 800 m boundary. This suggests that significant remineralization should have occurred at greater depths (even  $> 1000 \text{ m}$ ), and it is also reflected by the presence of  $Ba_{xs}$  maxima below 1000 m (see Fig. 4h and Appendix A). This was particularly salient when plotting  $Ba_{xs}$  contents vs. depths over the 27-day observation period (Fig. 8). The high, deep water  $Ba_{xs}$  values in Fig. 8 were not taken into account when integrating TNS-6 and E-1 profiles between 150 and 400 or even 150 and 800 m (Fig. 5e). Considering that the seafloor in the meander area is at about 2000 m depth, it seems unlikely that these high  $Ba_{xs}$  contents at depths  $> 1000 \text{ m}$  were due to sediment resuspension. Also, particle spectra for these sites do not reveal any bottom resuspension (Jouandet et al., 2014; Lasbleiz et al., 2014; van der Merve et al., 2015). Therefore, the high, deep ( $> 1000 \text{ m}$ )  $Ba_{xs}$  contents at TNS-6 and E-1 most likely reflects the fact that here significant remineralization of POC material actually did occur in the bathypelagic domain and even down to the seafloor. Note that suspended particles in the depth range containing the deep  $Ba_{xs}$  maxima were dominated by the  $< 2 \mu\text{m}$  size fraction (M. Zhou, personal communication, 2014). When integrating the  $Ba_{xs}$  contents from 150 m to the seafloor at stations TNS-6 and E-1, MR fluxes increase to 156 and  $184 \text{ mg C m}^{-2} \text{ day}^{-1}$ , respectively. Such C fluxes were similar to the EP values (maximum value of  $130 \text{ mg C m}^{-2} \text{ day}^{-1}$  at E-3) and suggested that the exported

POC was entirely remineralized in the water column leaving no C for transfer to the sediments.

Overall, the temporal pattern of mesopelagic remineralization described above reflects two successive events of particle transfer: a first transfer from a previous bloom (occurred before visiting TNS-6 and enduring at E-1) and a second transfer from E-4E to E-5. The first transfer was evidenced by the downward (to the bottom) propagation of the mesopelagic  $Ba_{xs}$  maximum signal, which mostly weakens at E-2. The second event was reflected by the occurrence again of important mesopelagic  $Ba_{xs}$  build-up at E-4E and E-5. Overall, our results indicated the large capacity of the polar front meander to transfer POC material to depth, but in contrast to station A3 on the plateau, this transfer was coupled to intense and near-complete POC remineralization (as also observed at E-4W and F-L). Between-site changes in mesopelagic carbon remineralization due to unequal biomass productivity and iron fertilization over the Kerguelen Plateau were thus relatively complex. Furthermore, the conditions in the meander area seems to corroborate results obtained in the iron-replete Subantarctic Zone east of the Tasman Plateau (Australian sector of the Southern Ocean; SAZ-Sense cruise; Jacquet et al., 2011a, b), where the mesopelagic remineralization efficiency reported was relatively high (on average 91 %) and the deep (> 600 m) carbon transfer weak (< 10 %). Finally, the important  $Ba_{xs}$  contents reported between 1000 and 2000 m during the first stages of the meander time-series support recent results indicating for the Southern Ocean that 1000 m is insufficient as an ocean-wide reference for carbon transfer and sequestration potential (Robinson et al., 2014).

## 5 Conclusions

Based on spatially and temporally well-resolved mesopelagic excess particulate Ba inventories, this work estimated mesopelagic POC remineralization above the Kerguelen Plateau and inside a permanent meander of the polar front to the east of plateau. The observed variability of mesopelagic remineralization reflects differences in the fate of the biomass that is exported to the deep ocean between the plateau and polar front. Results also reveal the patchiness of the seasonal advancement and of the establishment of remineralization processes between these sites. Our results indicate that a few days to weeks before the start of the cruise, the reference station R-2 experienced an export event that was efficiently remineralized in the upper mesopelagic layer. In terms of deep ocean carbon transfer efficiency, our results highlight that above the plateau (A3 site), mesopelagic remineralization is not a major barrier to organic matter transfer to the seafloor, with carbon transfer beyond 400 m reaching up to 87 % of EP during KEOPS 2, while in the polar front meander, remineralization of exported organic carbon in the upper 400 m is more efficient than above the plateau. In the meander area, remineralization may even balance export when including its effect in the deeper waters (up to 800 m and even deeper), thus resulting in a near-zero carbon transfer to sediment. A similar condition is also observed for sites at the margin of the plateau (E-4W) and the polar front (F-L).



## Appendix A

**Table A1.** Excess particulate biogenic Ba ( $Ba_{XS}$ ; pM) and particulate Al (nM) during KEOPS 2.

Station A3								Station RK2							
A3-1 CTD4				A3-2 CTD 107				R-2 CTD17				R-2 CTD 20			
Niskin no.	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]
23	11	224	35	23	11	122	16	24	21	110	107	17	356	546	12
21	42	217	64	21	40	140	12	23	40	0	693	15	507	239	17
19	104	345	65	19	81	141	10	22	80	95	49	13	609	226	7
17	152	234	19	17	126	82	27	20	100	131	27	11	812	267	3
15	173	244	19	15	151	119	14	18	126	168	5	10	1011	189	2
13	204	333	17	13	176	199	9	16	151	205	4	8	1520	201	4
11	227	235	8	11	202	186	14	14	203	334	3	6	1832	184	2
9	253	315	6	9	277	323	24	13	253	616	6	1	2473	286	3
7	279	480	8	7	303	359	32	12	304	834	16				
5	354	216	21	5	405	247	19	10	404	573	9				
1	474	1108	155	1	513	1842	186	9	507	430	10				
								7	608	367	4				
								5	708	337	10				
								1	911	368	13				

Station E															
E-1 CTD 27				E-1 CTD 30				E-2 CTD 43				E-3 CTD 50			
Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]
24	21	896	166	17	303	424	30	23	11	192	43	24	11	129	45
23	41	221	131	16	353	195	25	21	41	93	152	23	42	117	93
22	81	190	161	15	455	143	6	18	102	143	17	22	71	130	28
20	101	172	102	13	505	268	6	16	153	215	57	20	102	160	22
18	125	150	10	11	636	343	7	14	204	408	9	18	125	201	11
16	151	290	9	10	808	138	2	12	254	311	6	16	153	225	18
14	182	375	5	8	1011	442	4	10	305	227	4	14	203	193	3
13	253	450	12	6	1498	436	9	8	406	353	5	13	252	210	2
12	303	402	9	1	2042	326	7	7	507	371	9	12	304	309	6
10	403	327	10					6	609	271	10	10	404	316	7
9	505	230	6					5	813	297	14	9	505	419	64
7	605	305	10					4	1016	350	35	7	606	320	14
5	707	298	10					1	2020	302	11	5	707	193	12
1	913	309	7									1	912	265	5

E-3 CTD 55				E-4W CTD 81				E-4W CTD 87				E-4E CTD 94			
Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]
17	404	185	5	24	10	101	16	17	304	277	9	24	20	116	32
16	455	272	5	23	41	134	17	16	354	350	10	23	51	260	11
15	505	176	6	22	70	152	5	15	453	233	7	22	93	563	223
13	605	378	5	20	94	86	18	13	606	182	9	20	103	170	5
11	810	258	3	18	126	84	7	11	811	186	5	18	126	215	9
10	910	172	2	16	153	193	8	10	910	187	7	16	152	210	6
8	1012	184	3	14	203	312	4	8	1011	268	61	14	181	247	7
6	1214	228	6	13	252	628	17	6	1214	249	29	13	253	547	4
1	1908	237	9	12	304	488	12	1	1384	250	30	12	305	403	78
				10	406	594	11					10	404	408	26
				9	507	272	11					9	505	403	26
				7	607	418	12					7	606	298	13
				5	708	338	21					5	706	285	8
				1	909	294	14					1	912	245	65

Station E (continued)															
E-4E CTD 97				E-5 CTD 113				E-5 CTD 114							
Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]	Niskin	Depth [m]	$Ba_{XS}$ [pM]	Al [nM]				
21	404	199	2	10	911	111	3	24	11	210	5				
18	505	242	3	8	1011	266	16	23	41	196	14				
13	706	175	1	6	1214	256	15	22	82	245	4				
8	1012	212	11	1	1922	208	5	20	102	264	14				
7	1265	189	2					18	126	131	6				
6	1518	149	2					16	152	153	5				
5	1827	225	43					14	202	181	2				
4	2027	212	12					13	252	469	6				
1	2212	254	9					12	302	606	9				
								10	404	377	13				
								9	507	422	11				
								7	606	425	7				
								5	707	281	12				
								1	910	281	6				

Table A1. Continued.

West–east transect								Station F-L							
TEW-3 CTD38				TEW-8 CTD 47				F-L CTD 63				F-L CTD 68			
Niskin	Depth [m]	Ba <sub>xs</sub> [pM]	Al [nM]	Niskin	Depth [m]	Ba <sub>xs</sub> [pM]	Al [nM]	Niskin	Depth [m]	Ba <sub>xs</sub> [pM]	Al [nM]	Niskin	Depth [m]	Ba <sub>xs</sub> [pM]	Al [nM]
23	16	133	40	23	10	196	31	24	11	146	37	17	405	264	6
21	41	107	112	21	41	0	251	23	35	97	41	16	456	233	9
19	61	209	45	18	102	92	41	22	61	0	228	15	506	339	12
17	76	148	20	16	152	169	45	20	82	141	36	13	607	265	3
13	112	128	13	14	202	134	9	18	101	134	5	11	910	718	7
11	183	235	8	12	254	164	5	16	126	185	5	10	1013	257	5
9	253	391	11	10	304	268	12	14	151	221	5	8	1215	316	8
7	277	348	9	8	405	217	5	13	202	280	7	6	1772	225	7
5	404	356	8	7	507	319	5	12	252	399	8	1	2741	2999	131
1	545	242	13	6	609	209	8	10	303	420	7				
				5	809	330	14	9	404	305	26				
				4	1011	334	22	7	506	408	7				
				1	2812	11179	826	5	707	247	10				
								1	911	282	11				

North–south transect											
TNS-1 CTD15				TNS-6 CTD 10				TNS-8 CTD8			
Niskin	Depth [m]	Ba <sub>xs</sub> [pM]	Al [nM]	Niskin	Depth [m]	Ba <sub>xs</sub> [pM]	Al [nM]	Niskin	Depth [m]	Ba <sub>xs</sub> [pM]	Al [nM]
23	11	262	62	23	35	182	26	23	12	478	45
21	41	30	90	21	42	141	12	21	41	258	53
18	102	93	15	18	103	143	14	18	102	303	32
16	153	225	17	16	184	413	11	16	150	1008	33
14	202	289	5	14	204	461	8	14	205	341	10
12	253	228	2	12	255	298	5	12	254	447	6
10	304	521	4	10	306	505	7	10	305	481	4
8	405	346	3	8	407	474	4	8	405	312	3
7	506	230	1	7	509	464	9	7	505	208	3
6	607	352	13	6	610	315	15	6	606	283	7
5	809	234	4	5	813	269	9	5	707	325	11
4	1520	127	6	4	1526	362	13	4	910	376	35
1	2282	211	19	1	1886	410	21	1	1000	294	28

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